

## Carbon isotope fluctuations and Phosphorus accumulation during Late Devonian anoxic events in Erfoud Morocco

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The Late Devonian Frasnian-Famennian biotic crisis marks one of the most intriguing and largest mass extinctions in Earth history. The most severe biotic crises are the lower and upper Kellwasser events (LKWE-UKWE), which occurred at the end of the Frasnian and decimated the reef system, shallow benthos and some pelagic swimmers. These crises were associated with major sea-level rises, warm climate and widespread ocean anoxia.

Environmental and depositional changes across the Late Devonian Kellwassers events in the Anti-Atlas, near Erfoud, Morocco, are examined based on microfacies, mineralogy,  $\delta^{13}\text{C}$  values and phosphorus (P<sub>tot</sub>) analyses. Correlation based on  $\delta^{13}\text{C}_{\text{org}}$  with other sections (e.g. Benner, Germany, Baisha, China) reveals the Erfoud section as stratigraphically complete across the Frasnian-Famennian transition. Microfacies analyses suggest outer shelf to hemipelagic conditions. Both LKWE and UKWE correspond to laminated black shales deposited during a sea-level rise under anoxic conditions and coincide with increased detrital inputs (quartz, feldspars) reflecting intense weathering on land. However, P<sub>tot</sub> concentrations are quite different in the two Kellwasser horizons. The LKWE is almost completely depleted in P<sub>tot</sub>, contrary to the UKWE, which is characterized by more variable contents, ranging from very high contents (3500 ppm) to background values (<200ppm). This may indicate that the two Kellwasser levels reflect different environmental conditions. The LKWE appears to be more anoxic preventing the effective retention of P into the sediments. In contrast, the UKWE reflects more variable conditions with episodic oxic intervals.

## A multilayered ocean in the Ediacaran Yangtze platform? Insights from carbonate and organic matter paired $\delta^{13}\text{C}$

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The large carbon isotope fluctuations systematically reported for Ediacaran carbonate rocks are interpreted as evidence of strong environmental variations. New paired  $\delta^{13}\text{C}$  data on carbonates and their associated organic matter are presented for the Yangjiaping section of the Doushantuo Formation (Hunan, South China), which, integrated with previously reported ones [1, 2], allow the reconstruction of lateral variations of  $\delta^{13}\text{C}_{\text{carb}}$ ,  $\delta^{13}\text{C}_{\text{org}}$  and  $\Delta^{13}\text{C}_{\text{carb-org}}$  for a shelf-to-basin cross section of the Yangtze platform. Strong lateral heterogeneities are revealed, with complex variations of  $\delta^{13}\text{C}_{\text{carb}}$  and  $\Delta^{13}\text{C}_{\text{dol-org}}$  ( $\Delta^{13}\text{C}_{\text{dol-org}} = \delta^{13}\text{C}_{\text{carb}} - \delta^{13}\text{C}_{\text{org}}$ ) in the inner shelf section, phased variations (positive  $\delta^{13}\text{C}_{\text{carb}}$  and  $\Delta^{13}\text{C}_{\text{dol-org}}$  close to 29‰) in the shelf margin section, and negative  $\delta^{13}\text{C}_{\text{carb}}$  ( $\Delta^{13}\text{C}_{\text{dol-org}}$  as low as 20‰) in the basin.

Assuming that carbonate  $\delta^{13}\text{C}$  is acquired in bottom waters, we show that the spatial and temporal  $\delta^{13}\text{C}_{\text{carb}}$  and  $\Delta^{13}\text{C}_{\text{dol-org}}$  variations are compatible with a three-layered water column: (i) an oxic surface layer, which dissolved inorganic carbon (DIC) is probably in isotope equilibrium with the atmosphere; (ii) an intermediate euxinic layer with a DIC enriched in  $^{12}\text{C}$  due to organic matter oxidation by sulphate reduction; (iii) a deepest euxinic layer that seems to be restricted to the inner shelf lagoonal facies, lacking sulphate, and with a DIC enriched in  $^{13}\text{C}$  by methanogenesis. This model implies that some Ediacaran basins may have contained distinct DIC reservoirs, thus complicating our understanding of the global carbon cycle at the time.

[1] Guo *et al.* (2007) *Pal. Pal. Pal.* **254**, 140-157.

[2] McFadden *et al.* (2008) *PNAS* **105**, 3197-3202.